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# A Data Assimilation Perspective on the Initialization of Climate Prediction

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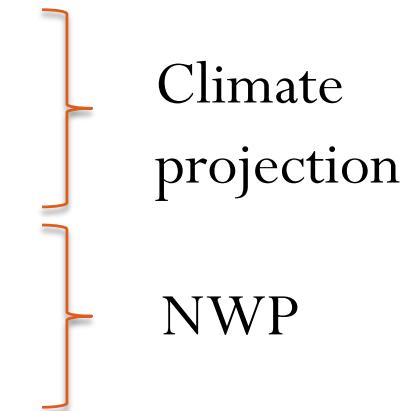
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# Backdrop: s2d prediction

s2d prediction:

- ✓ External forcing
- ✓ Coupled models
- ✓ Signature of initial condition



Initialization of  
s2d prediction:

- Full Field Initialization (FFI)



- Anomaly Initialization (AI)



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# Research objectives

- Propose a unified formalism from which the different initialization approaches can be derived
- Compare FFI and AI for a range of different observational and model error scenarios using an idealized coupled dynamics
- Propose two advanced formulations of FFI/AI to improve the assimilation of observations, and reduce the bias
- Illustrate the impact of initialization using a state-of-the-art Earth System Model

# Posing the problem

Climate model:

$$\frac{d\vec{x}}{dt} = F(\vec{x}, \vec{\lambda})$$

## Assumptions

Nature:

$$\frac{d\vec{x}^{nat}}{dt} = F(\vec{x}^{nat}, \vec{\lambda}^{nat}) + G(\vec{x}^{nat}, \vec{\lambda}^{nat})$$

$$\vec{x}, \vec{x}^{nat} \in X$$

$$[\vec{x}^{nat}] = [\vec{x}] = I$$

Observations:

$$\vec{y}_i^o = \vec{y}^o(t_i)$$

$$[\vec{\lambda}^{nat}] = [\vec{\lambda}] = P$$

$$t_i = i\tau, \quad i = 0, 1, \dots;$$

$$|G| \ll |F|$$

$$\vec{y}^o = \Phi(\vec{x}^{nat}) + \vec{\varepsilon}^o$$

$$[\vec{y}] = M \ll I$$

# DA formulation of standard initialization methods



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FFI: Model state is replaced by the best-possible available estimate of the actual state

$$\vec{x}^a = \vec{x}^b + H^T [\vec{y}^o - H\vec{x}^b] \quad [H] = M \times I$$

$\vec{x}^b$  : Background state from long “control” run of the model

AI: Assimilation of observed climate anomalies on an estimate of the model climate

$$\vec{y}^{pso} = \vec{y}^o - (\vec{y}^o - H\vec{x})$$

$$\vec{x}^a = \vec{x}^b + H^T [\vec{y}^{pso} - H\vec{x}^b]$$

# Least Square Initialization (LSI)

Using a least-square framework, the background and observations are linearly combined in order to minimize the expected analysis error covariance. Errors are assumed to be Gaussian.

$$\vec{x}^a = \vec{x}^b + BH^T [HBH^T + R]^{-1} [\vec{y}^o - H\vec{x}^b]$$

How do we estimate **B** to initialize s2d predictions?

Proxy:  $B^m = \overline{\alpha(\vec{x} - \bar{\vec{x}})(\vec{x} - \bar{\vec{x}})^T}$  Long time average

$$\alpha : \begin{cases} 1) & \text{Compensate estimation of the nature covariance} \\ 2) & \text{Optimize weighting} \end{cases}$$

# Exploring Parameter Uncertainty (EPU)

**Goal:** Reduce model drift caused by parametric error

**Approach:** Online bias correction method based on a short-time, linear approximation of the bias evolution

$$\frac{d\delta\vec{x}}{dt} \approx \left. \frac{\partial F}{\partial \vec{x}} \right|_{\vec{x}, \vec{\lambda}} \delta\vec{x} + \left. \frac{\partial F}{\partial \vec{\lambda}} \right|_{\vec{x}, \vec{\lambda}} \delta\vec{\lambda} \quad \rightarrow \quad \delta\vec{x} \approx \mathbf{M}_{t, t_0} \delta\vec{x}_0 + \int_{t_0}^t \mathbf{M}_{t, \tau} \delta\mu(\tau) d\tau$$

Deterministic!

Bias estimation  
proposal:

$$\begin{aligned} \vec{b}(t) &= \langle \delta\vec{x}(t) \rangle \\ &\approx \langle \delta\mu_0 \rangle [t - t_0] = \left. \langle \frac{\partial F}{\partial \vec{\lambda}} \rangle_{\vec{x}, \vec{\lambda}} \delta\vec{\lambda} \right| [t - t_0] \end{aligned}$$

Unknown!

# Exploring Parameter Uncertainty

Working  
Hypothesis:

**Identification of a set of uncertain parameters**

Identification of a range of possible parameter values  $[\lambda_{\min}, \lambda_{\max}]$

$$\begin{aligned}\vec{x}^{un}(t_i) &= \vec{x}(t_i) - \vec{b}(t_i) \\ &= \vec{x}(t_i) - \frac{\partial F}{\partial \lambda} \Bigg|_{\vec{x}(t_{i-1}), \vec{\lambda}} \delta \lambda_i \Delta T_{Bias} \quad t_i = i \Delta T_{Bias} \quad i = 1, 2, \dots\end{aligned}$$

$$\delta \vec{\lambda}_i \in \begin{cases} U(0, \vec{\lambda}^{\max} - \vec{\lambda}) & \text{if } \vec{\lambda} > \vec{\lambda} \\ U(\vec{\lambda}^{\min} - \vec{\lambda}, 0) & \text{if } \vec{\lambda} < \vec{\lambda} \end{cases}$$

# Idealized coupled model

Extratropical  
Atmosphere

Tropical  
Atmosphere

Tropical  
Ocean

$$\dot{x}_e = \sigma(y_e - x_e) - c_e(Sx_t + k_1)$$

$$\dot{y}_e = rx_e - y_e - x_e z_e + c_e(Sy_t + k_1)$$

$$\dot{z}_e = x_e y_e - bz_e$$

$$\dot{x}_t = \sigma(y_t - x_t) - c(SX + k_2) - c_e(Sx_e + k_1)$$

$$\dot{y}_t = rx_t - y_t - x_t z_t + c(SY + k_2) + c_e(Sy_e + k_1)$$

$$\dot{z}_t = x_t y_t - bz_t + c_z Z$$

$$\dot{X} = \tau\sigma(Y - X) - c(x_t + k_2)$$

$$\dot{Y} = \tau r X - \tau Y - \tau SXZ + c(y_t + k_2)$$

$$\dot{Z} = \tau SXY - \tau bZ - c_z z_t$$

Coupling

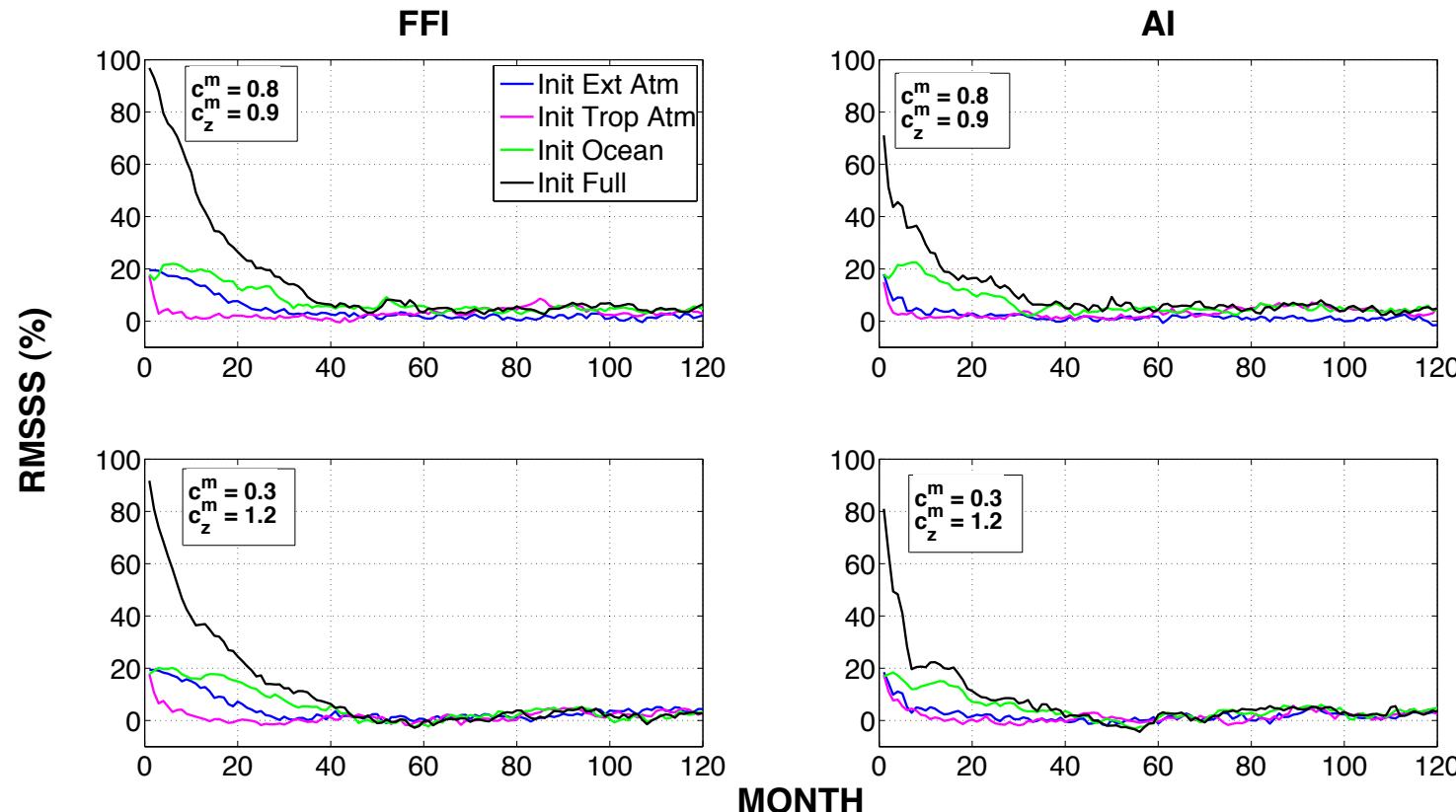
$c_e$

$c, c_z$

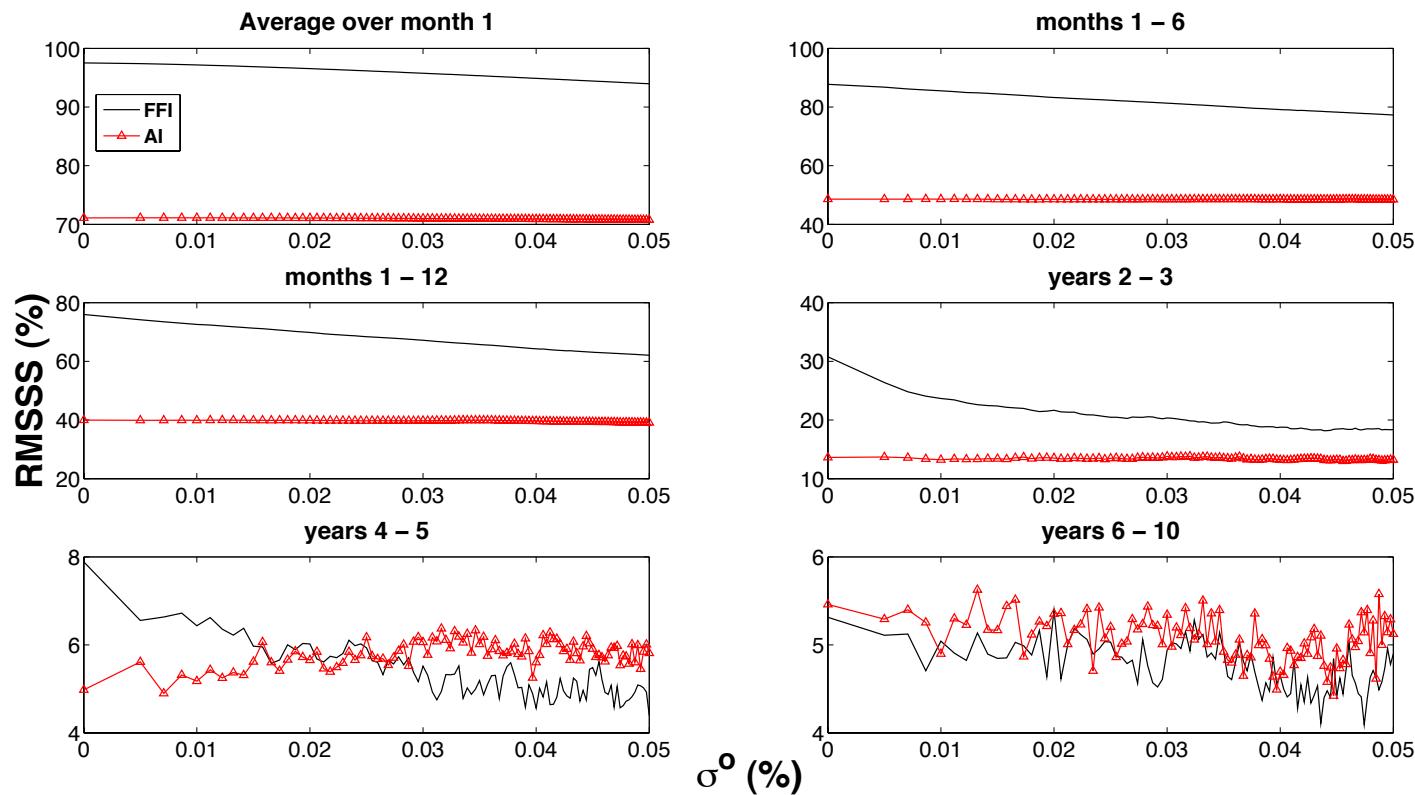
# Experimental Setup

- Nature: Solution to model equations with coupling strength  $c_e = 0.08$ ,  $c = c_z = 1$
- Introduction of parametric model error by modifying simultaneously  $c^m$ ,  $c_z^m$  in the range of (0.1 – 1.5)
- The observations are simulated by sampling the nature trajectory each month and adding a Gaussian white noise
- 10-yr long predictions are initialized each month for 30 yrs:  
Total of 360 startdates

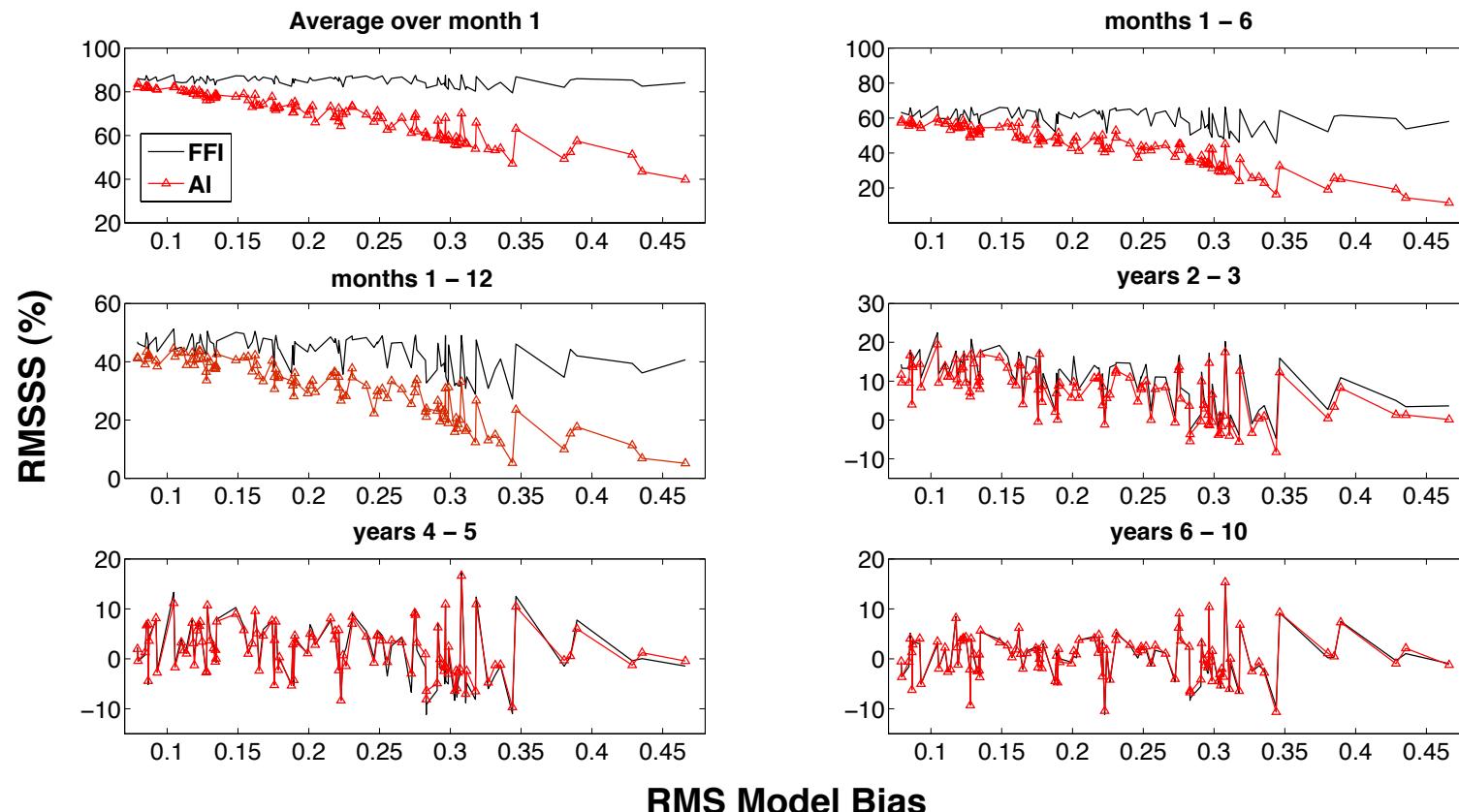
# Comparison AI / FFI



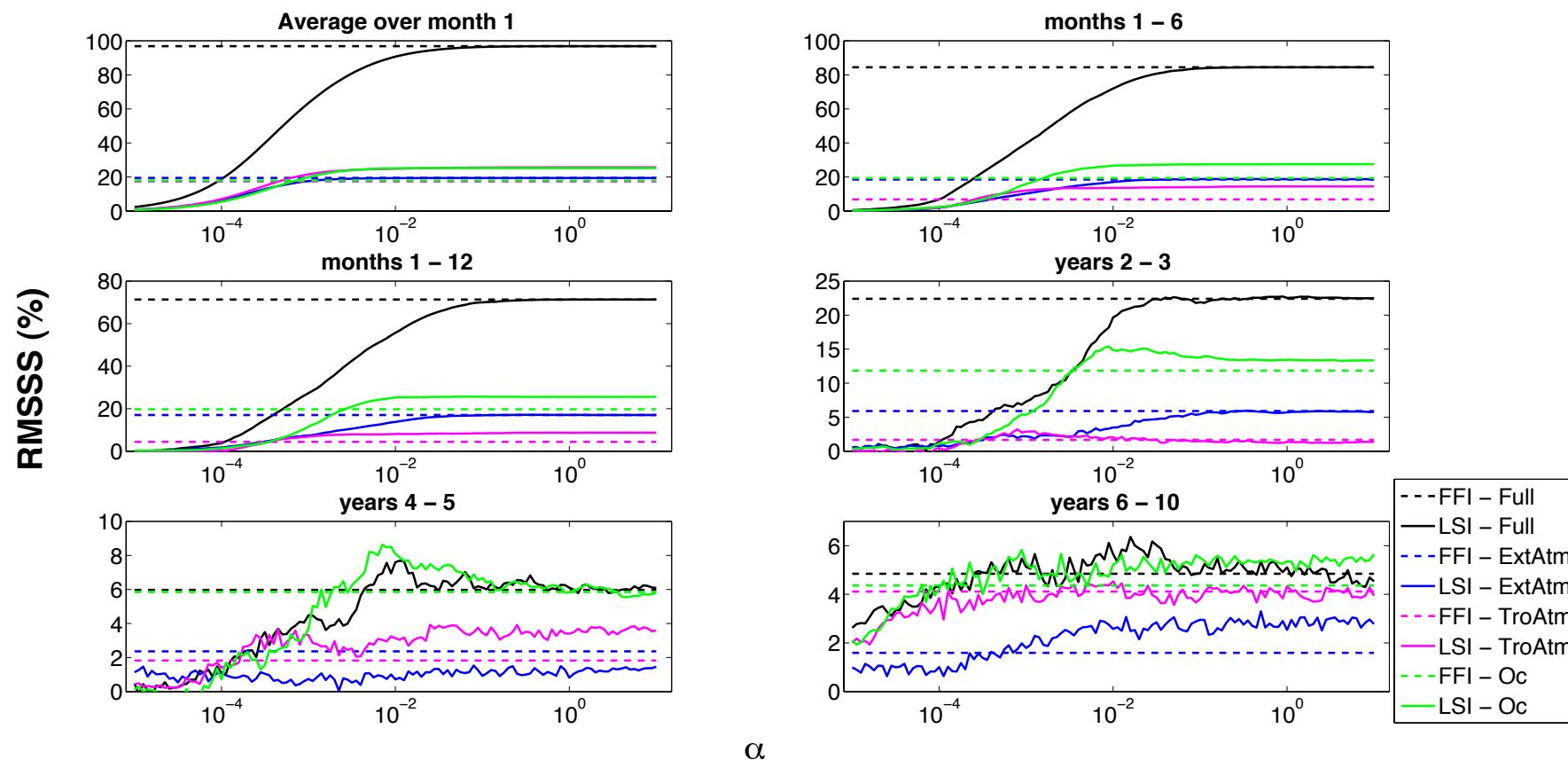
# Influence of observational error



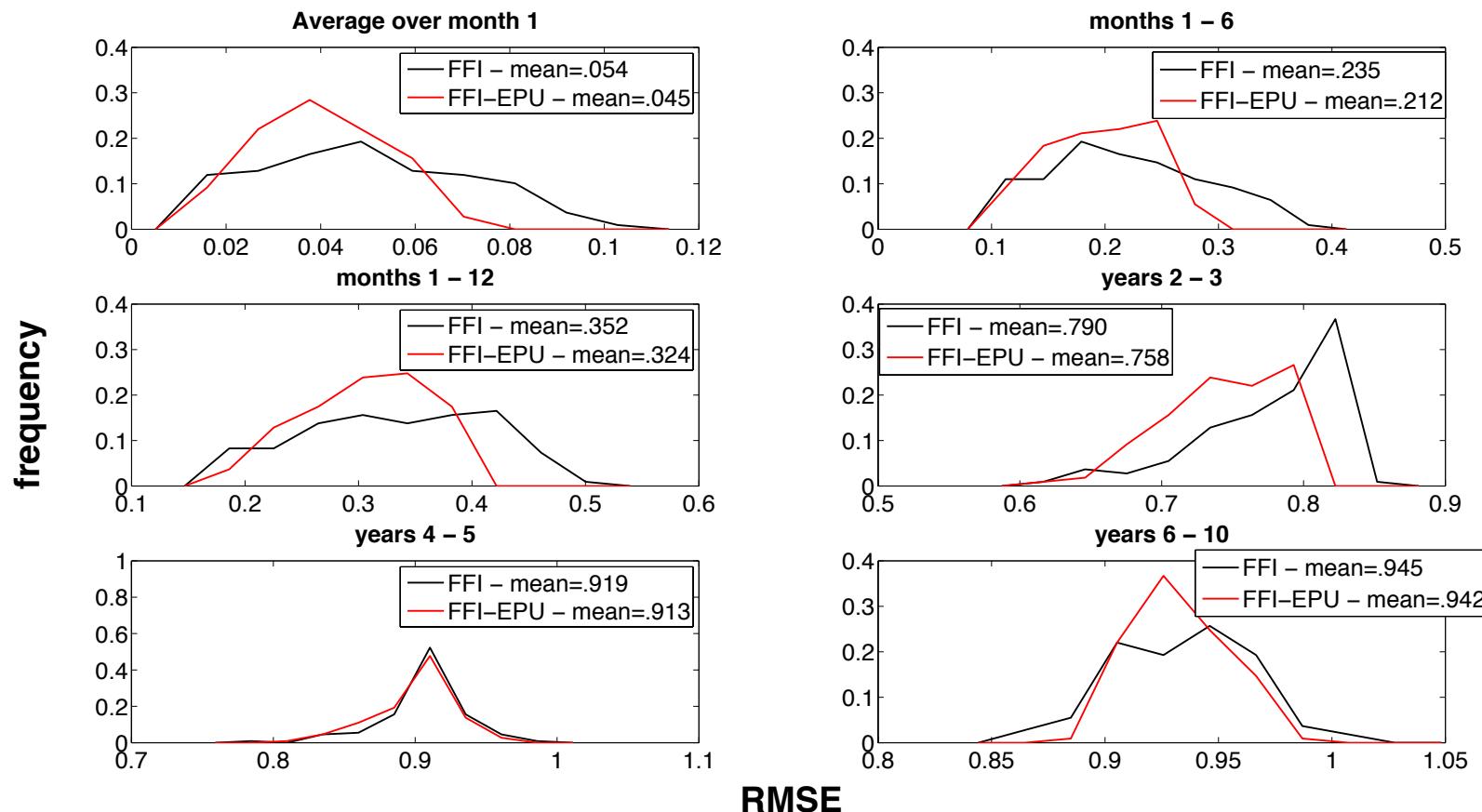
# Influence of model bias



# FF - LSI



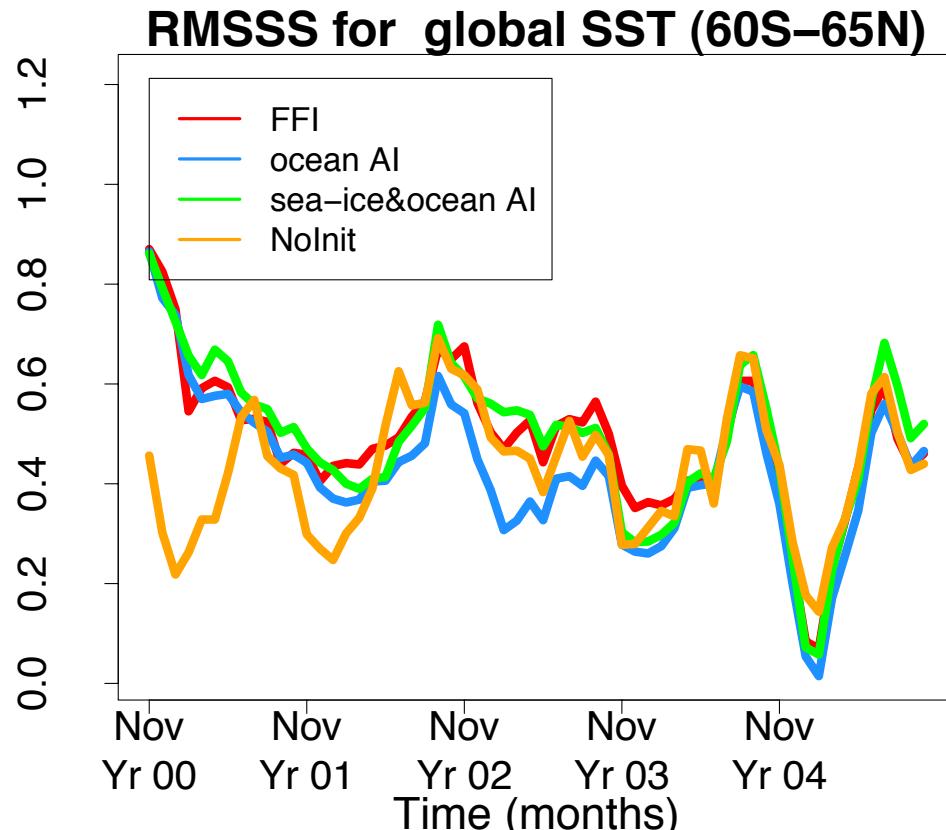
# FFI - EPU



# Coupled General Circulation Model: EC-Earth



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# Conclusions

- FFI slightly outperforms AI
- The ocean alone mainly determines the system's predictability horizon
- FFI sensitive to observational error, AI to model bias
- LSI helps improve prediction skill mainly by spreading out information, otherwise restricted to the observation subspace, to the entire model space
- Clear benefit of FFI-EPU up to 3 years in the forecast (4-17%)
- EC-Earth: added value from initialization for first forecast year, with only minor differences between initialization procedures

**Reference:** Carrassi A., R. Weber, V. Guemas, F. Doblas-Reyes, M. Asif, and D. Volpi 2013  
(submitted to Nonlin. Processes Geophys.)



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*Thank you*

2013 Summer School: Data Assimilation in Geosciences,  
CSCAMM