Calculating Storm Surge and Other Coastal Hazards Using Geoclaw

Kyle T. Mandli

Department of Applied Mathematics University of Washington Seattle, WA, USA

Modeling and Computations of Shallow-Water Coastal Flows, University of Maryland, 2010-10-14

1 Single-Layer Storm Surge Modeling

- GeoClaw
- Storm Surge Modeling with GeoClaw

2 Multi-Layer Storm Surge Modeling

- Multi-Layer Equations
- GeoClaw and Multiple Layers

David George: Mendenhall postdoctoral Fellow at the USGS Marsha Berger: Courant Institute, NYU Randy LeVeque: Applied Mathematics, University of Washington

Supported in part by NSF, ONR

1 Single-Layer Storm Surge Modeling

- GeoClaw
- Storm Surge Modeling with GeoClaw

2 Multi-Layer Storm Surge Modeling

- Multi-Layer Equations
- GeoClaw and Multiple Layers



- Wave-propagation class of high-resolution finite volume methods using a Godunov type scheme
- Available at www.clawpack.org/geoclaw
- Basic computation involves solving the Riemann problem at each cell interface



- Wave-propagation class of high-resolution finite volume methods using a Godunov type scheme
- Available at www.clawpack.org/geoclaw
- Basic computation involves solving the Riemann problem at each cell interface
- Ourrently includes:
 - 2d library for depth-averaged flows over topography
 - Well-balanced Riemann solvers that handle dry cells
 - General tools for dealing with multiple data sets at different resolutions
 - Tools for specifying regions where refinement is desired
 - Graphics routines (Matlab transitioning to Python)
 - Output of time series at gauge locations or on fixed grids

- USGS parameterization of fault zone
- Okada model

- USGS parameterization of fault zone
- Okada model

Refinement:

Coarsest level = 2°

- USGS parameterization of fault zone
- Okada model

Refinement:

- Coarsest level $= 2^{\circ}$
- Level 1 ightarrow Level 2, factor of 4 (30 minutes)

- USGS parameterization of fault zone
- Okada model

Refinement:

- Coarsest level = 2°
- Level 1 \rightarrow Level 2, factor of 4 (30 minutes)
- Level 2 \rightarrow Level 3, factor of 5 (6 minutes)

- USGS parameterization of fault zone
- Okada model

Refinement:

- Coarsest level = 2°
- Level 1 ightarrow Level 2, factor of 4 (30 minutes)
- Level 2 \rightarrow Level 3, factor of 5 (6 minutes)
- Δt Adaptive, based on CFL condition of grid





Kyle Mandli (UW Applied Math)







Kyle Mandli (UW Applied Math)











Chile Tsunami 2010: Dart Buoy Comparison



Chile Tsunami 2010: Continental Shelf



Chile Tsunami 2010: Continental Shelf

Storm Surge Modeling with GeoClaw

Storm Surge Modeling with GeoClaw

Method:

Assumed wind field

- Assumed wind field
- Add wind source term to momentum equation: $C_f \rho_{air} |W|^2$

- Assumed wind field
- Add wind source term to momentum equation: $C_f \rho_{air} |W|^2$
- C_f is a piece-wise defined, limited friction coefficient

- Assumed wind field
- Add wind source term to momentum equation: $C_f \rho_{air} |W|^2$
- C_f is a piece-wise defined, limited friction coefficient
- Treated using a source term splitting method

- Assumed wind field
- Add wind source term to momentum equation: $C_f \rho_{air} |W|^2$
- C_f is a piece-wise defined, limited friction coefficient
- Treated using a source term splitting method

Adaptive Refinement

• Currents are primarily used for the refinement criterion

Hurricane Forced Ocean Basin: Currents

Hurricane Forced Ocean Basin: Surface

Kyle Mandli (UW Applied Math)

Hurricane Forced Ocean Basin: Currents (Deep)

Hurricane Forced Ocean Basin: Surface (Deep)

Single-Layer Storm Surge Modeling

- GeoClaw
- Storm Surge Modeling with GeoClaw

2 Multi-Layer Storm Surge Modeling

- Multi-Layer Equations
- GeoClaw and Multiple Layers

Beyond Shallow Water Storm Surge Modeling












Storm surge model with two-layers:

• Use two layers - boundary layer and abyssal layer



Storm surge model with two-layers:

- Use two layers boundary layer and abyssal layer
- Wind only forces top layer



Storm surge model with two-layers:

- Use two layers boundary layer and abyssal layer
- Wind only forces top layer
- Use thermocline as boundary between layers

Multi-Layer Equations

Motivation: Provide more vertical structure

Multi-Layer Equations

Motivation: Provide more vertical structure

Integrate to an intermediate interface



• Internal wave speeds are much slower than gravity wave speeds

- Internal wave speeds are much slower than gravity wave speeds
- Governed strongly by ratio of densities

- Internal wave speeds are much slower than gravity wave speeds
- Governed strongly by ratio of densities
- Small surface waves can be accompanied by large internal waves

- Internal wave speeds are much slower than gravity wave speeds
- Governed strongly by ratio of densities
- Small surface waves can be accompanied by large internal waves
- Kelvin-Helmholtz instabilities

Multi-Layer Depth Integration

$$P = \rho_2 g h_2 + \rho_1 g (\eta_1 - z)$$
 $r = \rho_2 / \rho_1$

$$P = \rho_2 g h_2 + \rho_1 g (\eta_1 - z)$$
 $r = \rho_2 / \rho_1$

$$\int_{b}^{\eta_{1}} (u_{t} + (u^{2})_{x} + (uw)_{z}) dz = -\int_{b}^{\eta_{1}} P_{x} / \rho dz$$

$$P = \rho_2 g h_2 + \rho_1 g (\eta_1 - z)$$
 $r = \rho_2 / \rho_1$

$$\int_{b}^{\eta_{1}} (u_{t} + (u^{2})_{x} + (uw)_{z}) dz = -\int_{b}^{\eta_{1}} P_{x}/\rho dz$$

$$\frac{\partial}{\partial t} \int_{b}^{\eta_{1}} u dz + \int_{b}^{\eta_{1}} u^{2} dz = -\frac{1}{\rho_{1}} \int_{b}^{\eta_{1}} (\rho_{2}gh_{2} + \rho_{1}g(\eta_{1} - z))_{x} dz$$

$$P = \rho_2 g h_2 + \rho_1 g (\eta_1 - z)$$
 $r = \rho_2 / \rho_1$

$$\int_{b}^{\eta_{1}} (u_{t} + (u^{2})_{x} + (uw)_{z}) dz = -\int_{b}^{\eta_{1}} P_{x}/\rho dz \qquad \Rightarrow \\ \frac{\partial}{\partial t} \int_{b}^{\eta_{1}} u dz + \int_{b}^{\eta_{1}} u^{2} dz = -\frac{1}{\rho_{1}} \int_{b}^{\eta_{1}} (\rho_{2}gh_{2} + \rho_{1}g(\eta_{1} - z))_{x} dz \qquad \Rightarrow \\ (h_{1}u_{1})_{t} + (h_{1}u_{1}^{2})_{x} = -rgh_{1}(h_{2})_{x} - \frac{1}{2}gh_{1}^{2} - gh_{1}b_{x}$$

$$P = \rho_2 g h_2 + \rho_1 g (\eta_1 - z)$$
 $r = \rho_2 / \rho_1$

$$\begin{split} &\int_{b}^{\eta_{1}} (u_{t} + (u^{2})_{x} + (uw)_{z}) dz = -\int_{b}^{\eta_{1}} P_{x} / \rho dz \qquad \Rightarrow \\ &\frac{\partial}{\partial t} \int_{b}^{\eta_{1}} u dz + \int_{b}^{\eta_{1}} u^{2} dz = -1 / \rho_{1} \int_{b}^{\eta_{1}} (\rho_{2}gh_{2} + \rho_{1}g(\eta_{1} - z))_{x} dz \qquad \Rightarrow \\ &(h_{1}u_{1})_{t} + (h_{1}u_{1}^{2})_{x} = -rgh_{1}(h_{2})_{x} - 1 / 2gh_{1}^{2} - gh_{1}b_{x} \qquad \Rightarrow \\ &\frac{(h_{1}u_{1})_{t} + (h_{1}u_{1}^{2} + 1 / 2gh_{1}^{2})_{x} = -rgh_{1}(h_{2})_{x} - gh_{1}b_{x}}{(h_{1}u_{1})_{t} + (h_{1}u_{1}^{2} + 1 / 2gh_{1}^{2})_{x} = -rgh_{1}(h_{2})_{x} - gh_{1}b_{x}} \end{split}$$

Bottom
$$\begin{cases} (h_1)_t + (h_1u_1)_x = 0\\ (h_1u_1)_t + \left(h_1u_1^2 + \frac{1}{2}gh_1^2\right)_x = -gh_1(r(h_2)_x + b_x)\\ \\ Top \begin{cases} (h_2)_t + (h_2u_2)_x = 0\\ (h_2u_2)_t + \left(h_2u_2^2 + \frac{1}{2}gh_2^2\right)_x = -gh_2((h_1)_x + b_x) \end{cases}$$

Bottom
$$\begin{cases} (h_1)_t + (h_1u_1)_x = 0\\ (h_1u_1)_t + \left(h_1u_1^2 + \frac{1}{2}gh_1^2\right)_x = -gh_1(r(h_2)_x + b_x)\\ \\ Top \begin{cases} (h_2)_t + (h_2u_2)_x = 0\\ (h_2u_2)_t + \left(h_2u_2^2 + \frac{1}{2}gh_2^2\right)_x = -gh_2((h_1)_x + b_x) \end{cases}$$

Problem: Only conditionally hyperbolic

Bottom
$$\begin{cases} (h_1)_t + (h_1u_1)_x = 0\\ (h_1u_1)_t + \left(h_1u_1^2 + \frac{1}{2}gh_1^2\right)_x = -gh_1(r(h_2)_x + b_x)\\ \\ Top \begin{cases} (h_2)_t + (h_2u_2)_x = 0\\ (h_2u_2)_t + \left(h_2u_2^2 + \frac{1}{2}gh_2^2\right)_x = -gh_2((h_1)_x + b_x) \end{cases}$$

Problem: Only conditionally hyperbolic Write system as non-conservative, quasi-linear form

$$q_t + A(q)q_x = S(q)$$

Wave Speeds

$$((\lambda - u_1)^2 - gh_1)((\lambda - u_2)^2 - gh_2) - rg^2h_1h_2 = 0$$

$$((\lambda - u_1)^2 - gh_1)((\lambda - u_2)^2 - gh_2) - rg^2h_1h_2 = 0$$

Approximate wave speeds by assuming $|u_2 - u_1| \ll 1$ and $1 - r \ll 1$

$$((\lambda - u_1)^2 - gh_1)((\lambda - u_2)^2 - gh_2) - rg^2h_1h_2 = 0$$

Approximate wave speeds by assuming $|u_2 - u_1| \ll 1$ and $1 - r \ll 1$ External waves:

$$\lambda_{\text{ext}}^{\pm} = \frac{h_1 u_1 + h_2 u_2}{h_1 + h_2} \pm \sqrt{g(h_1 + h_2)}$$

$$((\lambda - u_1)^2 - gh_1)((\lambda - u_2)^2 - gh_2) - rg^2h_1h_2 = 0$$

Approximate wave speeds by assuming $|u_2 - u_1| \ll 1$ and $1 - r \ll 1$ External waves:

$$\lambda_{\text{ext}}^{\pm} = \frac{h_1 u_1 + h_2 u_2}{h_1 + h_2} \pm \sqrt{g(h_1 + h_2)}$$

ves:
$$g' = g(1-r)$$

$$\lambda_{ ext{int}}^{\pm} = rac{h_1 u_2 + h_2 u_1}{h_1 + h_2} \pm \sqrt{g' rac{h_1 h_2}{h_1 + h_2}} \left[1 - rac{(u_1 - u_2)^2}{g'(h_1 + h_2)}
ight]$$

Oscillating Wind Field

Augment multi-layer system with wind friction term in the top layer:

$$\operatorname{Top} \begin{cases} (h_2)_t + (h_2 u_2)_x = 0\\ (h_2 u_2)_t + \left(h_2 u_2^2 + \frac{1}{2}gh_2^2\right)_x = -gh_2((h_1)_x + b_x) + \tau_f |W|^2\\ \\ (h_1)_t + (h_1 u_1)_x = 0\\ (h_1 u_1)_t + \left(h_1 u_1^2 + \frac{1}{2}gh_1^2\right)_x = -gh_1(r(h_2)_x + b_x) \end{cases}$$

Modeling Considerations

Advantages:

Advantages:

• Vertical structure taken into account

Advantages:

- Vertical structure taken into account
- Modest increase in computational cost vs. 3D calculations

Advantages:

- Vertical structure taken into account
- Modest increase in computational cost vs. 3D calculations

Possible Difficulties:

Modeling Considerations

Advantages:

- Vertical structure taken into account
- Modest increase in computational cost vs. 3D calculations
- Possible Difficulties:
 - Hyperbolicity

Modeling Considerations

Advantages:

- Vertical structure taken into account
- Modest increase in computational cost vs. 3D calculations

Possible Difficulties:

• Hyperbolicity - Off of continental shelf, velocities should not violate hyperbolicity
Modeling Considerations

Advantages:

- Vertical structure taken into account
- Modest increase in computational cost vs. 3D calculations

- Hyperbolicity Off of continental shelf, velocities should not violate hyperbolicity
- Dry-state problem

Advantages:

- Vertical structure taken into account
- Modest increase in computational cost vs. 3D calculations

- Hyperbolicity Off of continental shelf, velocities should not violate hyperbolicity
- Dry-state problem Force bottom layer to become dry away from coastline

Advantages:

- Vertical structure taken into account
- Modest increase in computational cost vs. 3D calculations

- Hyperbolicity Off of continental shelf, velocities should not violate hyperbolicity
- Dry-state problem Force bottom layer to become dry away from coastline
- Computation of eigenvalues

Advantages:

- Vertical structure taken into account
- Modest increase in computational cost vs. 3D calculations

- Hyperbolicity Off of continental shelf, velocities should not violate hyperbolicity
- Dry-state problem Force bottom layer to become dry away from coastline
- Computation of eigenvalues State is near linear regime, approximations should be valid

Modeling Considerations: Hyperbolicity



Modeling Considerations: Dry-State Problem



Linearize about $u_1 = u_2 = 0$ and expand about 1 - r

Linearize about $u_1 = u_2 = 0$ and expand about 1 - rExternal waves:

$$\lambda_{\text{ext}}^{\pm} = \sqrt{g(h_1 + h_2)} - \frac{g^{1/2}h_1h_2}{2(h_1 + h_2)^{3/2}}(1 - r) + \mathcal{O}((1 - r)^2)$$

Linearize about $u_1 = u_2 = 0$ and expand about 1 - rExternal waves:

$$\lambda_{\text{ext}}^{\pm} = \sqrt{g(h_1 + h_2)} - \frac{g^{1/2}h_1h_2}{2(h_1 + h_2)^{3/2}}(1 - r) + \mathcal{O}((1 - r)^2)$$

Internal waves:

$$\lambda_{\rm int}^{\pm} = \sqrt{\frac{gh_1h_2}{h_1 + h_2}(1-r)} + \frac{g^{1/2}(h_1h_2)^{3/2}}{2(h_1 + h_2)^{5/2}}(1-r)^{3/2} + \mathcal{O}((1-r)^{5/2})$$

• Calculate linearized eigenvalues using only left and right states (no averaging)

- Calculate linearized eigenvalues using only left and right states (no averaging)
- Use f-wave approach to handle source terms

- Calculate linearized eigenvalues using only left and right states (no averaging)
- Use f-wave approach to handle source terms
 - Advantageous when problem is near steady state, $f(q)_{\star} \approx S(q)$

- Calculate linearized eigenvalues using only left and right states (no averaging)
- Use f-wave approach to handle source terms
 - Advantageous when problem is near steady state, $f(q)_{ imes} pprox S(q)$
- Refine based on speed of top layer and gradient of top surface and internal surface

Importance of the Steady State

Want to solve this problem:



Steady State Figures

Not this:



$$q_t + f(q)_x = S(q, q_x, ...)$$
 $q_t + A(q)q_x = S(q, q_x, ...)$

$$q_t + f(q)_x = S(q, q_x, ...)$$

Wave Propagation:

$$q_t + A(q)q_x = S(q, q_x, ...)$$

$$q_t + f(q)_x = S(q, q_x, ...)$$
 $q_t + A(q)q_x = S(q, q_x, ...)$

Wave Propagation:

() Compute eigenspace (speeds s and eigenvectors R) of our system

$$q_t + f(q)_x = S(q, q_x, ...)$$
 $q_t + A(q)q_x = S(q, q_x, ...)$

Wave Propagation:

- **(**) Compute eigenspace (speeds s and eigenvectors R) of our system
- **②** Compute jump in conserved quantities $q_i q_{i-1} = \delta$

$$q_t + f(q)_x = S(q, q_x, ...)$$
 $q_t + A(q)q_x = S(q, q_x, ...)$

Wave Propagation:

- **(**) Compute eigenspace (speeds s and eigenvectors R) of our system
- **2** Compute jump in conserved quantities $q_i q_{i-1} = \delta$

$$R\alpha = \delta$$

$$q_t + f(q)_x = S(q, q_x, ...)$$
 $q_t + A(q)q_x = S(q, q_x, ...)$

Wave Propagation:

- O Compute eigenspace (speeds s and eigenvectors R) of our system
- **2** Compute jump in conserved quantities $q_i q_{i-1} = \delta$

$$R\alpha = \delta$$

Construct waves and update grid cells based on wave speeds

$$\alpha^{p}r^{p} = \mathcal{W} \quad \Rightarrow \quad \sum \mathcal{W}s = \mathcal{A}^{\pm}\Delta q$$

$$q_t + f(q)_x = S(q, q_x, ...)$$
 $q_t + A(q)q_x = S(q, q_x, ...)$

Wave Propagation:

- O Compute eigenspace (speeds s and eigenvectors R) of our system
- **2** Compute jump in conserved quantities $q_i q_{i-1} = \delta$

$$R\alpha = \delta$$

Onstruct waves and update grid cells based on wave speeds

$$\alpha^{p}r^{p} = \mathcal{W} \quad \Rightarrow \quad \sum \mathcal{W}s = \mathcal{A}^{\pm}\Delta q$$

Solve source term
$$q_t = S(q)$$

 $q_t + f(q)_x = S(q, q_x, ...)$ $q_t + A(q)q_x = S(q, q_x, ...)$

F-Wave Propagation:

- Compute eigenspace (speeds s and eigenvectors R) of our system
- **3** Compute jump in conserved quantities $q_i q_{i-1} = \delta$

$$R\alpha = \delta$$

Onstruct waves and update grid cells based on wave speeds

$$lpha^{
m p} r^{
m p} = \mathcal{W} \quad \Rightarrow \quad \sum \mathcal{W} s = \mathcal{A}^{\pm} \Delta q$$

• Solve source term
$$q_t = S(q)$$

$$q_t + f(q)_x = S(q, q_x, ...)$$
 $q_t + A(q)q_x = S(q, q_x, ...)$

F-Wave Propagation:

- **(**) Compute eigenspace (speeds *s* and eigenvectors *R*) of our system
- **2** Compute jump in flux f(q) and add source term contributions

$$f(q_i) - f(q_{i-1}) - \Delta x S(q_{i-1/2}) = \Delta$$

$$R\alpha = \delta$$

O Construct waves and update grid cells based on wave speeds

$$\alpha^{p}r^{p} = \mathcal{W} \quad \Rightarrow \quad \sum \mathcal{W}s = \mathcal{A}^{\pm}\Delta q$$

• Solve source term
$$q_t = S(q)$$

$$q_t + f(q)_x = S(q, q_x, ...)$$
 $q_t + A(q)q_x = S(q, q_x, ...)$

F-Wave Propagation:

- **(**) Compute eigenspace (speeds s and eigenvectors R) of our system
- **Q** Compute jump in flux f(q) and add source term contributions

$$f(q_i) - f(q_{i-1}) - \Delta x S(q_{i-1/2}) = \Delta$$

O Project jump in conserved quantities △ onto eigenvector matrix to determine wave strengths

 $R\beta = \Delta$

Onstruct waves and update grid cells based on wave speeds

$$\alpha^{p}r^{p} = \mathcal{W} \quad \Rightarrow \quad \sum \mathcal{W}s = \mathcal{A}^{\pm}\Delta q$$

• Solve source term
$$q_t = S(q)$$

$$q_t + f(q)_x = S(q, q_x, ...)$$
 $q_t + A(q)q_x = S(q, q_x, ...)$

F-Wave Propagation:

- **(**) Compute eigenspace (speeds s and eigenvectors R) of our system
- **Q** Compute jump in flux f(q) and add source term contributions

$$f(q_i) - f(q_{i-1}) - \Delta x S(q_{i-1/2}) = \Delta$$

 Project jump in conserved quantities Δ onto eigenvector matrix to determine wave strengths

$$R\beta = \Delta$$

Onstruct waves and update grid cells based on wave speeds

$$\beta^{p}r^{p} = \mathcal{Z} \quad \Rightarrow \quad \sum \mathcal{Z} = \mathcal{A}^{\pm} \Delta q$$

Solve source term
$$q_t = S(q)$$

$$q_t + f(q)_x = S(q, q_x, ...)$$
 $q_t + A(q)q_x = S(q, q_x, ...)$

F-Wave Propagation:

- **(**) Compute eigenspace (speeds s and eigenvectors R) of our system
- **Q** Compute jump in flux f(q) and add source term contributions

$$f(q_i) - f(q_{i-1}) - \Delta x S(q_{i-1/2}) = \Delta$$

 Project jump in conserved quantities Δ onto eigenvector matrix to determine wave strengths

$$R\beta = \Delta$$

Onstruct waves and update grid cells based on wave speeds

$$\beta^{p}r^{p} = \mathcal{Z} \quad \Rightarrow \quad \sum \mathcal{Z} = \mathcal{A}^{\pm} \Delta q$$

• Solve source term $q_t = S(q)$

$$q_t + f(q)_x = S(q, q_x, ...)$$
 $q_t + A(q)q_x = S(q, q_x, ...)$

F-Wave Propagation:

() Compute eigenspace (speeds *s* and eigenvectors *R*) of our system

(2) Compute jump in flux f(q) and add source term contributions

$$f(q_i) - f(q_{i-1}) - \Delta x S(q_{i-1/2}) = \Delta$$

 Project jump in conserved quantities Δ onto eigenvector matrix to determine wave strengths

$$R\beta = \Delta$$

O Construct waves and update grid cells based on wave speeds

$$\beta^{p}r^{p} = \mathcal{Z} \quad \Rightarrow \quad \sum \mathcal{Z} = \mathcal{A}^{\pm} \Delta q$$

Hurricane Forced Ocean Basin: Top Surface

Kyle Mandli (UW Applied Math)

Hurricane Forced Ocean Basin: Currents

Hurricane Forced Ocean Basin: Depths

Hurricane Forced Ocean Basin: Top Surface

Hurricane Forced Ocean Basin: Currents

Hurricane Forced Ocean Basin: Depths

Conclusions
• Multi-layer equations provide increased model accuracy at modest computational overhead

- Multi-layer equations provide increased model accuracy at modest computational overhead
- Storm surge flows avoid many common problems with multi-layer equations

- Multi-layer equations provide increased model accuracy at modest computational overhead
- Storm surge flows avoid many common problems with multi-layer equations
- F-wave algorithm with a linearized approximation simplifies calculation and appears stable

- Multi-layer equations provide increased model accuracy at modest computational overhead
- Storm surge flows avoid many common problems with multi-layer equations
- F-wave algorithm with a linearized approximation simplifies calculation and appears stable

- Multi-layer equations provide increased model accuracy at modest computational overhead
- Storm surge flows avoid many common problems with multi-layer equations
- F-wave algorithm with a linearized approximation simplifies calculation and appears stable

Future Work:

• Fix what appears to be a "odd-even" decoupling in simulation

- Multi-layer equations provide increased model accuracy at modest computational overhead
- Storm surge flows avoid many common problems with multi-layer equations
- F-wave algorithm with a linearized approximation simplifies calculation and appears stable

- Fix what appears to be a "odd-even" decoupling in simulation
- Add a transverse solver

- Multi-layer equations provide increased model accuracy at modest computational overhead
- Storm surge flows avoid many common problems with multi-layer equations
- F-wave algorithm with a linearized approximation simplifies calculation and appears stable

- Fix what appears to be a "odd-even" decoupling in simulation
- Add a transverse solver
- Handle dry state for lower layer

- Multi-layer equations provide increased model accuracy at modest computational overhead
- Storm surge flows avoid many common problems with multi-layer equations
- F-wave algorithm with a linearized approximation simplifies calculation and appears stable

- Fix what appears to be a "odd-even" decoupling in simulation
- Add a transverse solver
- Handle dry state for lower layer
- Improve refinement criteria

- Multi-layer equations provide increased model accuracy at modest computational overhead
- Storm surge flows avoid many common problems with multi-layer equations
- F-wave algorithm with a linearized approximation simplifies calculation and appears stable

- Fix what appears to be a "odd-even" decoupling in simulation
- Add a transverse solver
- Handle dry state for lower layer
- Improve refinement criteria
- Viscous drag between layers

- Multi-layer equations provide increased model accuracy at modest computational overhead
- Storm surge flows avoid many common problems with multi-layer equations
- F-wave algorithm with a linearized approximation simplifies calculation and appears stable

- Fix what appears to be a "odd-even" decoupling in simulation
- Add a transverse solver
- Handle dry state for lower layer
- Improve refinement criteria
- Viscous drag between layers
- Real bathymetry

- Clawpack: www.clawpack.org
- GeoClaw: www.clawpack.org/geoclaw
- The GeoClaw software for depth-averaged flows with adaptive refinement, by M.J. Berger, D.L. George, R.J. LeVeque, and K.T. Mandli

www.clawpack.org/links/awr10/ or arXiv:1008.0455v1

• Paper with details in *Acta Numerica* in preparation, draft available soon.